#### Chapter Three Rainfall-Runoff Analysis

- Linear System Theory and Rainfall-Runoff Analysis
  - Unit hydrograph theory
    - From Stream flow Data
    - Synthetically
    - "Fitted" Distributions
  - Instantaneous unit hydrograph (IUH)
  - IUH analysis methods:
    - S-Hydrograph
    - Conceptual model
    - Fitting Harmonic analysis Fourier transforms
    - Theoretically from Laplace transforms
- River and Reservoir Flood Routing
  - Flood Routing

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- Reservoir flood routing methods:
- Linear Muskingum method:
- Multiple reach Muskingum method
- Nonlinear Muskingum method:



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# Unit Hydrograph Theory

- A mathematical concept
- Linear in nature
- Definition

The hydrograph that results from 1-unit of excess precipitation (or runoff) spread uniformly in space and time over a watershed for a given duration.

- The key points :
  - ✓1-unit of EXCESS precipitation
  - ✓ Spread uniformly over space evenly over the watershed
  - ✓ Uniformly in time the excess rate is constant over the time interval
  - ✓ There is a given duration







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# Unit Hydrograph "Lingo"

- Duration
- Lag Time
- Time of Concentration
- Rising Limb
- Recession Limb (falling limb)
- Peak Flow
- Time to Peak (rise time)
- Recession Curve
- Separation
- Base flow







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# **Methods of Developing UHG's**

- From Stream flow Data
  - Synthetically
    - Snyder
    - SCS
    - Time-Area (Clark, 1945)
  - "Fitted" Distributions





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# **Derived Unit Hydrograph**

#### **Rules of Thumb :**

 $\Sigma$ ... the storm should be fairly uniform in nature and the excess precipitation should be equally as uniform throughout the basin. This may require the initial conditions throughout the basin to be spatially similar.



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- Separate the base flow from direct runoff,
- Determine the volume of direct runoff, and
- Divide the ordinate of the direct runoff hydrograph by observed runoff depth.

$$r_d = \frac{\Delta t \sum_{i=1}^n q_i}{A}$$





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- The duration of the derived unit hydrograph is found by examining the precipitation for the event and determining that precipitation which is in excess.
- This is generally accomplished by plotting the precipitation in hyetograph form and drawing a horizontal line such that the precipitation above this line is equal to the depth of excess precipitation as previously determined.
- This horizontal line is generally referred to as the  $\phi$ -index and is based on the assumption of a constant or uniform infiltration rate.





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#### **Estimating Excess Precip.**





#### Example

Given the following hydrograph of a given watershed having drainage

area of 104 km<sup>2</sup> derive the unit hydrograph for the watershed

Date	Hour	Stream flow	Base flow
16-Feb	06:00	11	8
	08:00	170	8
	10:00	260	6
	12:00	266	6
	14:00	226	8
	16:00	188	9
	18:00	157	11
	20:00	130	12
	22:00	108	14
	24:00	91	16
17-Feb	02:00	76	17
	04:00	64	19
	06:00	54	21
	08:00	46	22
	10:00	38	24
	12:00	32	26
	14:00	27	27







# Synthetic UHG's

- Snyder
  - SCS
  - Time-area



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- Since peak flow and time of peak flow are two of the most important parameters characterizing a unit hydrograph, the Snyder method employs factors defining these parameters, which are then used in the synthesis of the unit graph (Snyder, 1938).
- The parameters are  $C_p$ , (peak flow factor), and  $C_t$ , (lag factor).
- The basic assumption in this method is that basins which have similar physiographic characteristics are located in the same area will have similar values of  $C_t$  and  $C_p$ .
- Therefore, the method is prefer to develop UH for un-gaged basins, which are near or similar to gaged basins for which these coefficients can be determined.



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# Snyder's Method

- Study area: US highlands. 10 ~ 10,000 mi<sup>2</sup> or (30 ~30,000 km<sup>2</sup>)
- Properties characterizes the response of watershed under various rainfall inputs:
  - (a) Lag time  $(t_L)$ ;
  - (b) Duration of rainfall excess (tr);
  - (c) Time base of UH (t<sub>b</sub>);
  - (d) Time to peak (t<sub>p</sub>);
  - (e) Peak discharge of UH (q<sub>p</sub>);
  - (f) Hydrograph time width at 50% and 75% (W50, W75) of peak flow



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### Snyder's Method

<u>Lag time</u> (t<sub>p</sub>): time from the center of rainfall – excess to the UH peak

 $t_p = C_t (L L_c)^{0.3}$ 

Where  $t_p = Time of lag [hrs];$ 

C<sub>t</sub> = Coefficient which is a function of watershed slope and shape, (~ 1.35 – 1.60, for steeper slope, C<sub>t</sub> is smaller); L = length of the main channel [km] or [mi];

 $L_c$  = length along the main channel to the point nearest to the watershed centroid





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#### Snyder's Method

<u>UH Duration  $(t_{\underline{r}})$ :</u>

$$t_{\rm r} = t_{\rm p} / 5.5$$

where  $t_r$  and  $t_p$  are in [hrs].

If the duration of UH is other than  $t_r$ , then the lag time needs to be adjusted as

 $t_{np} = t_p + 0.25 (t_R - t_r)$ 

where  $t_{np}$  = adjusted lag time;  $t_{R}$  = desired UH duration.

UH Peak Discharge (q\_p):  
$$q_p = \frac{2.78AC_p}{t_p}$$

where  $C_p = \text{coefficient}$  accounting for flood wave and storage condition, 0.56 ~ 0.69 as miller; 0.23 - 0.67 clark;  $q_p = \text{specific}$  discharge,  $[m^3/s]$ 



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#### Snyder's Method

Time Base 
$$(T_{\underline{b}})$$
:  $\mathbf{t}_{\underline{b}} = (72 + 3t_{\underline{p}}) \text{ or } \mathbf{T}_{\underline{b}} = 5(t_{\underline{np}} + 0.5 t_{\underline{r}})$ 

<u>UH Widths</u>:

$$W_{50} = \frac{5.87}{q_{pru}^{1.08}} \qquad \qquad W_{75} = \frac{3.354}{q_{pru}^{1.08}}$$

where

 $W_{50}$ , are in hours; Usually, 2/3 of the width after the peak of UH  $W_{75}$  are in hours; Usually, 1/3 of the width is before UH peak  $q_{pru}$  = the peak discharge of the UH per unit drainage area (Qp/A)

The final shape of the Snyder unit hydrograph is controlled by the equations for width at 50% and 75% of the peak of the UHG



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Two catchments A and B are considered meteorologically similar . Their

catchment characteristics are given below

Catchment A	Catchment B		
L= 30km	L= 45km		
L <sub>ca</sub> = 15km	L <sub>ca</sub> = 25km		
A= 250km <sup>2</sup>	A= 400km <sup>2</sup>		

For catchment A, a 2hr unit hydrograph was developed and was found to have a peak discharge of 50m3/sec. The time to peak from the beginning of the rainfall excess in this unit hydrograph was 9.0hr. Using the Snyder's method develop a unit hydrograph for catchment B.





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#### SCS Dimensionless UH

	Time Ratios	Discharge Ratios	Mass Curve Ratios	_
L 'I	(t/t <sub>p</sub> )	$(q/q_p)$	$(Q_a/Q)$	_
	0	.000	.000	_
	.1	.030	.001	
	.2	.100	.006	
	.3	.190	.012	
	.4	.310	.035	
	.5	.470	.065	
	.6	.660	.107	
	.7	.820	.163	
	.8	.930	.228	
	.9	.990	.300	
	1.0	1.000	.375	
	1.1	.990	.450	
	1.2	.930	.522	
	1.3	.860	.589	B h - UH at time t
	1.4	.780	.650	7 Vs = Accumulated volume at time t 7 V = Total volume
	1.5	.680	.700	.6 t = A selected time tp = time from beginning of tp = time trom beginning of
	1.6	.560	.751	δ 5- ( / )
	1.7	.460	.790	
	1.8	.390	.822	a- / <sup>#</sup> / -c
	1.9	.330	.849	2- / /
	2.0	.280	.871	.1
	2.2	.207	.908	
	2.4	.147	.934	
	2.6	.107	.953	Figure 10.4 Dimensionless unit nyorograph and mass curve (after Soll Conservation Service, 1971).
	2.8	.077	.967	
	3.0	.055	.977	
	3.2	.040	.984	
	3.4	.029	.989	



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### **Triangular Representation**

 $T_{b} = 2.67 \text{ x } T_{p}$ 

$$T_{r} = T_{b} - T_{p} = 1.67 \text{ x } T_{p}$$

$$Q = \frac{q_{p}T_{p}}{2} + \frac{q_{p}T_{r}}{2} = \frac{q_{p}}{2}(T_{p} + T_{r})$$

$$q_p = \frac{2Q}{T_p + T_r}$$

$$q_{p} = \frac{654.33 \times 2 \times A \times Q}{T_{p} + T_{r}}$$

$$q_p = \frac{484 \, A \, Q}{T_p}$$

The 645.33 is the conversion used for delivering 1inch of runoff (the area under the unit hydrograph)

from 1-square mile in 1-hour (3600 seconds).





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## **Triangular Representation**

Comes from the initial assumption that 3/8 of the volume under the UHG is under the rising limb and the remaining 5/8 is under the recession limb.

<b>General Description</b>	<b>Peaking Factor</b>	Limb Ratio (Recession to Rising)
Urban areas; steep slopes	575	1.25
Typical SCS	484	1.67
Mixed urban/rural	400	2.25
Rural, rolling hills	300	3.33
Rural, slight slopes	200	5.5
Rural, very flat	100	12.0

Dura	ation and times	
$T_p = \frac{D}{2} + L$	$T_{c} + D = 1.7 T_{p}$	Easting tion and a second
		For estimation purposes :
L = 0.6 * T	D	$D = 0.133 T_c$
$L = 0.0$ $I_c$	$\frac{-}{2}$ + 0.6 $T_c = T_p$	



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# **Hypothetical Example**

- A 190 mi<sup>2</sup> watershed is divided into 8 isochrones of travel time.
- The linear reservoir routing coefficient, R, estimated as 5.5 hours.
- A time interval of 2.0 hours will be used for the computations.

2

Map Area #	Bounding Isochrones	Area (mi <sup>2</sup> )	Cumulative Area (mi <sup>2</sup> )	Cumulative Time (hrs)
1	0-1	5	5	1.0
2	1-2	9	14	2.0
3	2-3	23	37	3.0
4	3-4	19	58	4.0
5	4-5	27	85	5.0
6	5-6	26	111	6.0
7	6-7	39	150	7.0
8	7-8	40	190	8.0
TOTAL		190	190	8.0

#### **Rule of Thumb**

R - The linear reservoir routing coefficient can be estimated as approximately 0.75 times the time of concentration.

7





Time (hrs) 

#### **Incremental Area**



Synthetic time-area curve -The U.S. Army Corps of **Engineers (HEC 1990)** 

for  $(0.5 \le \text{Ti} \le 1.0)$ 



#### **General guiding Rule in synthetic UH**

- The UH is better calculated from strong and isolated storm events... Not many of them happened in the last 120 days
- Verification:
  - Find The watershed with the perfect storms and storm information
  - Create a synthetic Hydrograph from the UH and test it for other stream gages in the watershed
  - Compare the peak flows from the UH and the rationale method (C i A)





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#### **Fitting with Probability distribution**

#### **Fitting a Gamma Distribution**

$$f(t;a,b) = \frac{t^{a}e^{-t/b}}{b^{a+1}\Gamma(a+1)}$$





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## Project in UH

- Take one river gauging station from the country gauging stations dataset.
- Develop UH for the gauging station based on gauged data
- Develop the catchment characteristics
- Develop UH synthetic UH for that catchment
- Compare the peak of UH and give your opinion on the results



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